

Visualized effects of oxidation and temperature on pseudo-single-domain Fe₃O₄ particles examined by environmental TEM and off-axis electron holography

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Magnetic minerals in rocks record the direction and intensity of the ambient magnetic field during formation, providing, for example, varied information of the geomagnetic field and past tectonic plate motions. Magnetic carriers in the single domain (SD) grain size range (< 100 nm) are regarded as ideal paleomagnetic recorders because of their strong remanence and high magnetic stability, with potential relaxation times that are greater than the age of the Universe. However, magnetic signals from rocks are often dominated by small multi-domain grains that exhibit magnetic recording fidelities similar to those of SD grains (termed pseudo-SD (PSD)). To reliably interpret paleomagnetic signals from these PSD grains, mechanisms which induce and alter their magnetic remanence must be fully understood.

Magnetite (Fe₃O₄) is the most magnetic naturally occurring mineral on Earth, carrying the dominant magnetic signature in rocks and providing a critical tool in paleomagnetism [1]. The oxidation of Fe₃O₄ to other iron oxides, such as maghemite (γ-Fe₂O₃) and hematite (α-Fe₂O₃), is of particular interest as it influences the preservation of remanence of the Earth's magnetic field by Fe₃O₄. Further, signals in Fe₃O₄ grains are acquired in the direction of the geomagnetic field as they cool below their Curie temperature (T_C) of ~ 580 °C. However, current understanding of the effects of oxidation and thermomagnetic behaviour of PSD remanence in Fe₃O₄ grains is restricted to bulk magnetic measurements and numerical models. Fortunately, the advanced transmission electron microscopy (TEM) techniques of off-axis electron holography and environmental TEM allow for magnetic imaging of nano-scale minerals during *in situ* heating under controlled atmospheres.

In the present study, synthetic Fe₃O₄ particles in the PSD size range (< 200 nm) were heated *in situ* in an environmental TEM to a temperature of 700 °C in 8 mbar of O₂ [2]. The oxidation of the Fe₃O₄ particles was monitored qualitatively using electron energy-loss spectroscopy (EELS). Fig. 1a shows a native smooth-surfaced Fe₃O₄ grain and complementary EELS analysis of the Fe 2p $L_{2,3}$ edge is in good agreement with that of pure Fe₃O₄ (Fig. 1b, blue spectra). Close examination of the EEL spectra after *in situ* heating revealed the development of pre-peaks (arrowed) close to the Fe 2p $L_{2,3}$ edge (Figs. 1b, red spectra) that are indicative of oxidation. The associated effect of oxidation on their magnetic remanence was investigated using electron holography, in the form of reconstructed magnetic induction maps, where the oxidised grains exhibited a change in remanent strength and direction (Fig. 1c,d). The thermomagnetic behaviour of Fe₃O₄ particles in the PSD size range is also investigated using electron holography (Fig. 2) [3-5]. Magnetic induction maps recorded during *in situ* heating to just below the T_C reveal the thermal stability of several PSD Fe₃O₄ grains by visualizing their vortex domain states. The vortex states in small Fe₃O₄ grains are shown to rotate close to its T_C (Fig. 2b-d), rather than remaining thermally stable as seen in the vortex states of larger Fe₃O₄ grains.

[1] L. Néel *Adv. Phys.* 4, 191-242 L (1995).

[2] T. P. Almeida *et al.*, *Nat. Commun.* 5, 5154 (2014).

[3] T. P. Almeida *et al.*, *Sci. Adv.*, 2, e1501801 (2016).

[4] T. P. Almeida *et al.*, *Geophys. Res. Lett.*, 41, 7041 - 7047 (2014).

[5] T. P. Almeida *et al.*, *Geophys. Res. Lett.*, 43, 8426-8434 (2016).

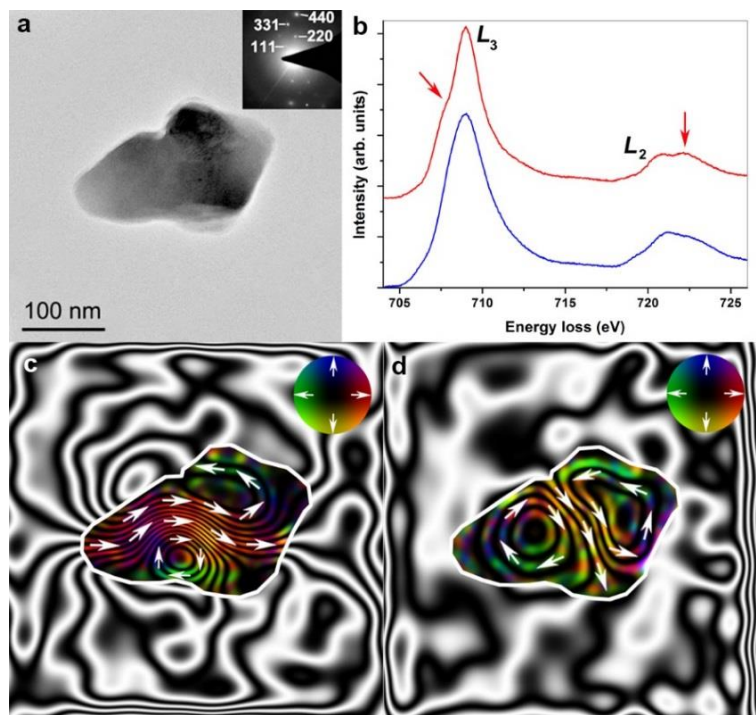


Fig. 1 (a) TEM image of an Fe_3O_4 particle and (b) Fe edges of the EEL spectrum acquired before (blue) and after oxidation (red). (c,d) Magnetic induction maps of the particle (c) before; and (d) after oxidation.

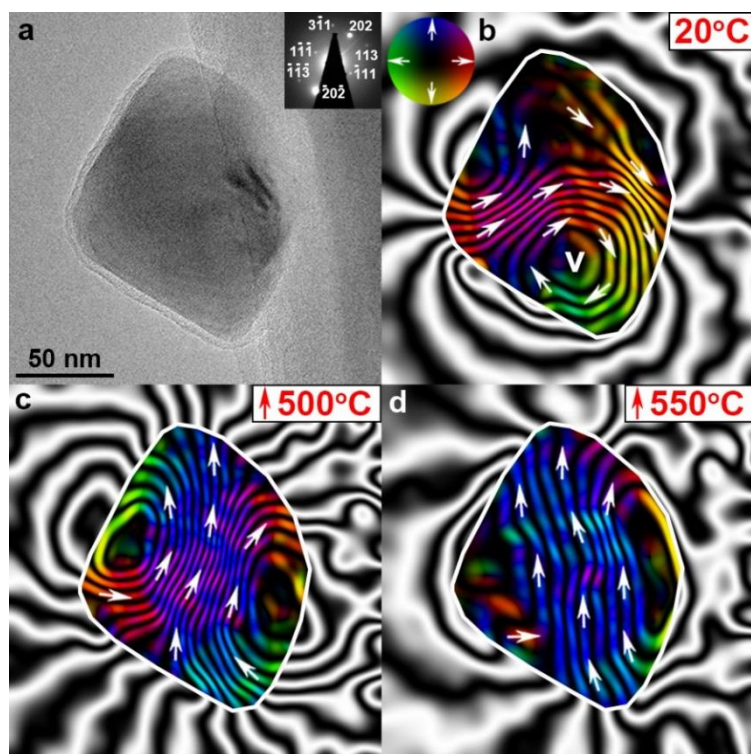


Fig. 2 (a) TEM image of an Fe_3O_4 particle and magnetic induction maps taken (b) at 20°C ; during *in situ* heating to (c) 500°C ; and (d) 550°C .